

# TECTONIC EVOLUTION OF THE CHUGACH-PRINCE WILLIAM TERRANE: SHUMAGIN ISLANDS AND KENAI PENINSULA, ALASKA

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## INTRODUCTION

An outstanding question in the assembly of the North American Cordillera is the accretion and subsequent translation history of the Campanian to Eocene Chugach-Prince William terrane (CPW) that extends for ~2200 kilometers along southern Alaskan margin (Fig. 1). The CPW is a thick accretionary complex primarily composed of deep-water turbidites with abundant quartzofeldspathic and volcanic-lithic sandstones and basaltic rocks (Plafker, 1998) that was intruded by near-trench plutons of the Sanak-Baranof belt (Hudson et al., 1979) inferred to be related to the subduction of the Kula-Farallon or Kula-Resurrection plate ridge (Bradley et al., 2003; Haeussler et al. 2003; Cowan, 2003). There are two prevailing hypotheses for the position of formation of the CPW along the Cordilleran margin: 1) The CPW formed more or less in place and ridge subduction progressed from west to east along the southern Alaskan margin during subduction of the Resurrection plate (Haeussler et al., 2003); or 2) the CPW formed far to the south and was intruded by near-trench plutons at 48-49°N during coastwise translation of the CPW (Cowan, 2003). The possible formation of the CPW far to the south and subsequent translation along the continental margin in the Paleocene and Eocene may be a defining event in Cordilleran tectonics and makes testable predictions for the provenance and thermal evolution of these rocks.

This contribution summarizes our findings from the second year of a three-year Keck Geology Consortium and NSF-supported project to understand the tectonic and thermal evolution of the Chugach-Prince William Terrane in southern Alaska. In summer 2012 our

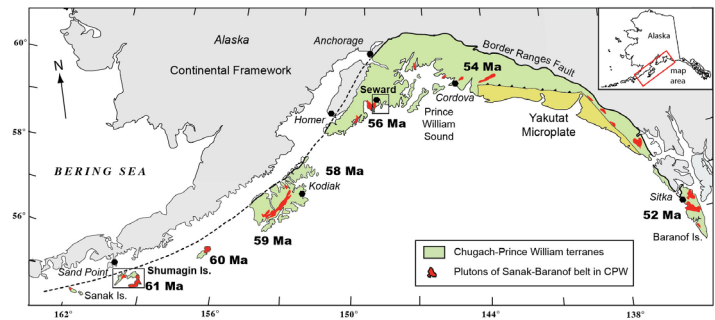
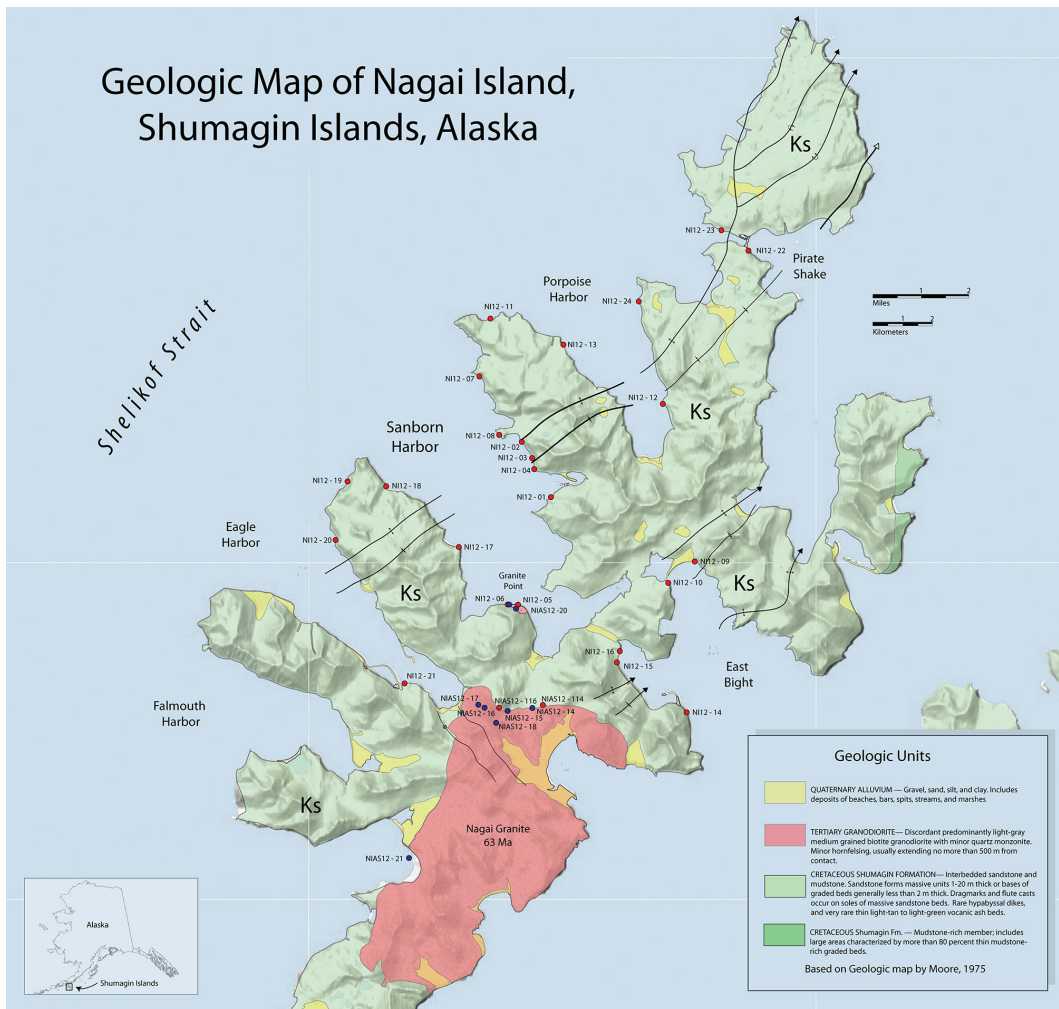


Figure 1. Extent of the Chugach-Prince William terrane in south central Alaska. 2012 field areas are indicated by black boxes. Pluton locations and dates are from Farris and Patterson (2009) and Bradley et al. (2003). Map modified from Pavlis and Roeske (2007).

research group included five students from five different undergraduate institutions who spent four weeks in the field working on the Kenai Peninsula, Resurrection Bay near Seward, and the Shumagin Islands, near Sand Point, Alaska (Fig. 2).

## PROJECT GOALS

The specific objectives for this project (summer 2012) were to: 1) determine the depositional setting and source for the Shumagin Formation exposed on Nagai Island in the Shumagin Archipelago near Sand Point, Alaska; 2) determine the crystallization age and geochemistry of the Shumagin Batholith complex, part of the Paleocene Sanak-Baranof belt of plutons; and 3) define the thermal history of the Shumagin Formation and batholith complex using zircon fission track dating. In addition to these objectives focused on the western exposures of the CPW in the Shumagin Islands, we were interested in: 4) determining the



*Figure 2. Geologic map of the Shumagin Islands, Alaska. Sample locations shown as red dots (Shumagin Formation), and blue dots (Nagai pluton). Map modified from Moore (1974).*

origin and depositional age of CPW flysch exposed stratigraphically above the Resurrection Ophiolite near Seward; and finally, 5) measure the Hf isotope composition of detrital zircon from across the CPW to help define provenance. We complement much of these efforts with ongoing studies of zircon crystallinity using Raman spectroscopy.

## RESEARCH

### Age, Provenance, and Thermal History of the Shumagin Formation

**Carly Roe** (Lawrence University) measured 1053 detrital zircon U/Pb dates from ten samples of volcanic-lithic and arkosic sandstones from the Shumagin Formation on Nagai Island. She shows that the maximum depositional ages from Nagai Island are closely clustered and range from 73–77 Ma confirming a Late Cretaceous age for the Shumagin Formation. She also obtained a U/Pb zircon crystallization age of  $73.7 \pm 1.2$  Ma from an interbedded tuff further

confirming the Late Cretaceous age of these rocks. The major age populations from the Shumagin Formation do not appear to correlate with units of similar age immediately inboard of the Shumagin Islands (Kuskokwim Group). However, they do compare well with correlative units along strike including the Kodiak and Ghost Rocks Formations on Kodiak Island, Valdez and Orca Groups in Prince William Sound, and the Yakutat Formation near Yakutat, Alaska.

**Alex Short** (University of Minnesota-Morris) collected 12 samples from the Shumagin Batholith for major and trace element geochemistry and used U/Pb zircon dating from two samples to determine a crystallization age for the batholith. He obtained concordant crystallization ages of  $61.7 \pm 0.7$  Ma and  $62.6 \pm 0.7$  Ma, and both samples showed some inheritance. The geochemistry of the batholith is distinctly peraluminous with most of the samples containing magmatic muscovite, and one sample contains a large (~ 1 cm) xenocryst of cordierite. Alex uses these data and field relationships to suggest that the Shumagin

Batholith formed from basaltic underplating and partial melting of the CPW accretionary complex.

**Mike Deluca** (Union College) uses detrital zircon fission track (ZFT) dating of samples from the Shumagin Formation and shows that these rocks have a complex cooling history with samples near the plutons yielding: 1) thermally reset cooling ages of 44-53 Ma; 2) samples with partially reset ages that have a wide range of cooling ages, but with a young population at 45-58 Ma; and 3) not thermally reset samples with a wide range of cooling ages that reflect the cooling ages of the original source regions. He double dated (ZFT and U/Pb) one sample from the thermally reset group and one sample from the partially reset group and shows that zircons with high effective U concentrations are more likely to be reset than grains with low effective U concentrations. From these data, Deluca suggests that the Shumagin Formation was buried to depths at or near the lower bound of the zircon partial annealing zone, was intruded and heated locally by the Shumagin batholith at ~62 Ma, and after intrusion, the rocks remained at depth and cooled slowly into the Eocene (c. 50 Ma) before passing to higher crustal levels.

### **Age of Clastic Cover to the Resurrection Peninsula Ophiolite**

**Rose Pettiette** (Washington and Lee) takes on a controversy over the age and stratigraphic relationship of clastic strata that structurally overlies the Resurrection Ophiolite near Seward, Alaska. The 57 Ma Resurrection Ophiolite is mapped by some workers as a thrust slice into Upper Cretaceous Valdez Group flysch (e.g. Bradley and Miller, 2006); however, Kusky and Young (1999) suggest that the clastic rocks overlying the ophiolite are in depositional contact and therefore younger than, or coeval with, the crystallization age of the ophiolite. Pettiette reports new detrital zircon U/Pb dates from four samples (n=404) from Thumb Cove, Humpy Cove, and Nash Road, across the bay from Seward. One sample (RB12-04), collected at the end of Humpy Cove, is from a thin-bedded, medium-grained sandstone interbedded with (and cross-cut by) basaltic rocks of the Resurrection Ophiolite. The maximum depositional age of RB12-04 is 57 Ma given by a robust mode formed from the youngest four zircons. Pettiette also

shows that the clastic strata in Resurrection Bay are nearly identical in zircon grain age distribution to the Paleocene part of the Orca Group 70-80 km to the NE in Prince William Sound. This tie of the Resurrection Ophiolite to the flysch of the Prince William terrane is significant because it removes the stratigraphic uncertainty of paleomagnetic data obtained from the Resurrection Peninsula ophiolite that indicate a paleolatitude  $13 \pm 9^\circ$  south of the present location to near present day northern Washington (Bol et al., 1992).

### **Hafnium Isotope Composition of Zircon from the Chugach-Prince William Terrane**

**Nick Roberts** (Carleton College) presents several hundred integrated U/Pb ages and Hf isotope compositions from four widely separated locations along the arcuate belt of the CPW to reveal a range of source regions with dramatically different zircon crystallization and mantle separation ages. Zircons from the Shumagin Formation yield positive  $\epsilon_{\text{Hf}}(t)$  values consistent with partial melting of a relatively juvenile source region. This is in contrast to similar-aged zircons from the other three sampling areas (from west to east), Kodiak Island, Prince William Sound, and Yakutat, where  $\epsilon_{\text{Hf}}(t)$  values range from +11.9 to -26.5 suggesting that Phanerozoic zircons from these areas crystallized from melts derived from a heterogeneous source region that included juvenile and Precambrian crust. Taken together, the integrated U/Pb and Hf isotope data show that the origin and provenance of detrital zircons from the CPW varies systematically along strike. Phanerozoic U/Pb ages have a strong Coast Mountains batholith signature with more juvenile crust in the magmatic source region in the Shumagin Formation, compared to tectonostratigraphically equivalent rocks in Kodiak, PWS, and Yakutat.

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