

Synchronous barium peaks in high-resolution profiles of calcite and aragonite marine bivalve shells

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Abstract Barium/calcium profiles of bivalve shells are characterized by flat background signals periodically interrupted by sharp peaks, with the background signals correlated with water Ba/Ca. To test if the peaks are an environmental signal related to productivity, we analyzed high-resolution Ba/Ca profiles in bivalve shells that grew adjacent to one another. Two aragonitic *Saxidomus giganteus* show remarkable similarity over a decade of growth, clearly indicating an environmental forcing. Four calcitic *Pecten maximus* shells also record synchronous Ba/Ca peaks, again indicating an exogenous control. The Ba/Ca peaks, however, start ~40 days after the crash of a bloom,

while sedimentation takes place immediately following the bloom. Barite formation in settling phytoplankton flocs, as has been previously proposed, is clearly not the cause of these peaks. Other possible causes, such as dissolved Ba in ambient water, spawning, shell organic matter content, and kinetic growth rate effects are also discussed, but none provide satisfactory explanations. Background shell Ba partition coefficients ($Ba/Ca_{\text{carbonate}}/Ba/Ca_{\text{water}}$) for both the calcitic shells (0.18) and aragonitic shells (0.16) are similar to that previously reported for the calcitic *Mytilus edulis* (~0.1). We suggest that Ba/Ca peaks in bivalve shells are caused by an as yet undetermined environmental forcing, while background Ba/Ca levels are a good indication of dissolved Ba/Ca in the water; both are independent of shell mineralogy.

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Introduction

Elemental signatures in biogenic carbonates such as corals and foraminifers are well-known proxies of climate change and past environmental conditions (see Lea 2003 for review). However, elemental signatures in bivalve shells are generally not well correlated with environmental conditions, due to strong vital effects (e.g., Purton et al. 1999; Takesue and van Geen 2004; Gillikin et al. 2005a; Lorrain et al. 2005; Surge and Walker 2006; Carré et al. 2006; Freitas et al. 2006). On the other hand, Ba/Ca ratios do seem to have potential as an environmental proxy.

There is a nearly ubiquitous Ba/Ca signal in all bivalve shells: a more or less constant Ba/Ca background that is interrupted with sharp episodic peaks (Stecher et al. 1996; Vander Putten et al. 2000; Lazareth et al. 2003; Carré et al. 2006; Gillikin et al. 2006). The $Ba/Ca_{\text{background}}$ has been shown to be a reliable proxy of Ba/Ca_{water} in *Mytilus edulis*

(Gillikin et al. 2006), and the sharp peaks seem to be reproducible between different shells growing at the same location, indicating an environmental forcing (Vander Putten et al. 2000; Carré et al. 2006). Stecher et al. (1996) first noted these peaks in the aragonite shells of two marine clam species. They postulated that these peaks may represent periods of barite (BaSO_4) formation immediately following phytoplankton blooms. Barite is considered to form in organic-rich decaying phytoplankton flocs, where BaSO_4 saturation conditions might prevail (Dehairs et al. 1980; Bishop 1988). Vander Putten et al. (2000) suggested that the Ba/Ca peaks in calcite mussel shells provide excellent markers for spring phytoplankton blooms, but noted there was no direct relationship between peak height and chlorophyll *a* (chl *a*) concentrations. Lazareth et al. (2003) also postulated that the Ba/Ca peaks in a tropical calcitic bivalve were related to phytoplankton blooms.

Other potential causes of these peaks could be short increases of dissolved Ba, biological partitioning of Ba to calcification sites (such as during spawning), or differences in shell micro-structure (see Gillikin et al. 2006 for review). In bivalves, these peaks are not generally considered to be correlated with riverine discharge (see discussion below), in contrast to corals (e.g., McCulloch et al. 2003; Sinclair and McCulloch 2004).

The aim of this study is to compare Ba/Ca profiles between shells grown at the same location, and then to compare shell data with environmental parameters. We also look at shell background Ba/Ca data from these species to determine if they are potentially a proxy of water Ba/Ca, as Gillikin et al. (2006) proposed for *Mytilus edulis*. In order to determine if these Ba/Ca peaks are reproducible between specimens growing at the same locality, which would suggest an extrinsic cause, we sampled two aragonitic *Saxidomus giganteus* shells from Puget Sound (Washington, USA) that provide a 10 year Ba/Ca record, and four calcitic *Pecten maximus* shells from the Bay of Brest (France) that provide precisely time-resolved

Ba/Ca profiles. Several environmental parameters recorded during shell growth were also available from these sites, such as temperature, salinity, and chl *a* concentrations.

Materials and methods

We chose *Saxidomus giganteus* (Fig. 1a) and *Pecten maximus* (Fig. 1b) for analysis for several reasons. *Saxidomus giganteus* provides (1) a long-term record (~10 years) from a single shell allowing ontogenic effects on peaks and background Ba/Ca ratios to be determined, and (2) an example of an aragonitic bivalve. The calcitic *Pecten maximus*, although providing only a short record (<1 year), allows precise matching between shell and environmental variables. Shells of both of these species have been studied extensively (e.g., Chauvaud et al. 1998, 2005; Lorrain et al. 2000, 2005; De Ridder et al. 2004; Gillikin 2005; Gillikin et al. 2005a,b; Freitas et al. 2006; Kingston et al. 2008).

The two *S. giganteus* shells were collected from intertidal sediments of Puget Sound (Washington, USA), and have been described in detail elsewhere (Gillikin et al. 2005a, b). Shells were sampled in cross-section along the axis of maximum growth, thus providing a high-resolution elemental profile spanning ~10 years. To compare data between shells and environment, the $\delta^{18}\text{O}$ profiles were used (presented in Gillikin et al. 2005b). Briefly, the $\delta^{18}\text{O}$ signal was fit between the two shells (*x*-axis of one shell scaled to the other). Then, the $\delta^{18}\text{O}$ signal was converted to a temperature (SST) signature, using a constant $\delta^{18}\text{O}_{\text{water}}$ and the paleotemperature equation of Böhm et al. (2000). This modeled temperature was then fit to environmental SST (scaling the *x*-axis only) using the algorithm described in De Ridder et al. (2004; see Gillikin et al. 2005b for complete details); this resulted in a coefficient of determination (R^2) of ~0.8. Considering we did not account for variability in water $\delta^{18}\text{O}$, this is a fairly good correlation, but nevertheless does

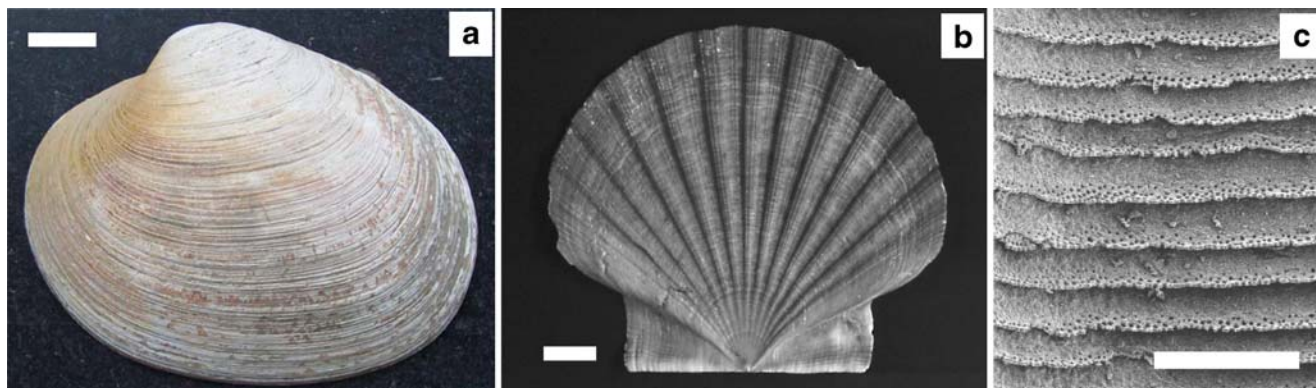


Fig. 1 Pictures of the bivalve species used in this study: *Saxidomus giganteus* (a, scale bar=1 cm) and *Pecten maximus* (b, scale bar=

1 cm). Details of the daily striae on *P. maximus* shells are also shown (c, scale bar=1 mm)

lead to an unknown error in determining the exact timing of the samples taken from the shell.

Shell Ba/Ca ratios were determined using laser-ablation ICP-MS (Fisons-VG frequency quadrupled Nd-YAG laser (266 nm) coupled to a Fisons-VG PlasmaQuad II+ICP-MS) as described in Gillikin et al. (2005a, 2006). Data from LA-ICP-MS were calibrated using both the NIST 610 (values from Pearce et al. 1997) and the USGS MACS1 (values from S. Wilson, USGS, unpublished data, 2004). The laser was shot directly in the drill holes used to generate the $\delta^{18}\text{O}$ data, allowing a precise alignment between the two methods (see Gillikin et al. 2005a, 2006). Environmental SST and chl *a* data are from the Washington State Department of Ecology (<http://www.ecy.wa.gov/>).

The *P. maximus* shells (three age class I, with one winter growth check, and one age class II, with two winter growth checks) were collected from 30 m depth in the Bay of Brest (France) in 2003; the age class I shells are the same shells presented in Lorrain et al. (2005). Daily growth rate was determined by measuring distances between successive daily growth striae (external growth increments; Fig. 1c) as described in Chauvaud et al. (1998). The method used to assign a precise date to each increment has already been well developed in several papers (e.g., Lorrain et al. 2004; Chauvaud et al. 2005). Each year, the date of growth restart is calculated by counting and measuring the daily increments from the edge of the shell to the first winter mark on 50 juveniles from the same cohort sampled at different times during the growth season. Dates are thus assigned by backdating from the collection date. A mean profile is obtained by minimizing the sum of the differences between the different growth profiles (Walter and Pronzato 1994). This mean curve is then used as a reference to assign dates to the striae from any new specimen. A potential error can come from the fact that some individuals may experience a short growth hiatus of a couple of days as a result of a disturbance. An additional source of error may be that a few striae can be broken from manipulation at the ventral edge of the scallop. The error is on the magnitude of a few days—approximately 1 to 7 days, depending on the scallop. The excellent correlation ($R^2=0.92$) found between $\delta^{18}\text{O}$ and temperature in *P. maximus* from this same location (Chauvaud et al. 2005) confirms that the backdating is indeed very precise. In addition, a maximum error of 1 week would not change any of our interpretations.

Carbonate powders representing ~1–6 days growth (covering the 2003 growth season) were milled from the *P. maximus* shells, and acid digested before being analyzed on an HR-ICP-MS (Finnigan Element2) as described in Lorrain et al. (2005). Reproducibility of an in-house calcite bivalve shell reference material was 6.4% (Ba/Ca=1.21±0.08 $\mu\text{mol/mol}$). Water was continuously monitored over the lifetime of the scallops for SST and chl *a* (fluorescence),

by means of a Sea-bird 19 sonde. At the same location and in the same year, reproductive activity of 3-year-old scallops was assessed by fortnightly sampling of 20 individuals. Three-year-old individuals were used because their larger gonads allow more precise sampling; however, previous studies show high reproductive synchronism between age groups in this species (Paulet and Fifas 1989). For each date, a gonad index was calculated as the dry weight of the gonad divided by the shell weight, and multiplied by 1,000 (see Paulet et al. 1992 for details).

Results and discussion

Ba/Ca Peaks

The inter-individual synchronicity of peaks in the shells of the two species is clearly evident in Figs. 2, 3, and 4. The very large, sharp peaks in 1993 (two peaks), 1994, and 1999 are timed exactly in both *S. giganteus* shells (red arrows, Fig. 2; Fig. 3a). Ba/Ca was independently fit using $\delta^{18}\text{O}$ cycles (Gillikin et al. 2005b), so the exact peak match is strong evidence of a good time base reconstruction. However, the offset in the 1996 peaks (red circle, Fig. 2; Fig. 3b) could be due to improper $\delta^{18}\text{O}$ alignment between the two shells, as this was the most difficult section of the shells to date (see Gillikin et al. 2005b). This offset approximates 80 days, which is considered the maximum error on the fit between shells. However, it should be kept in mind that the other regions of the shell seem to be fit more precisely (e.g., Fig. 3a). The smaller peaks (<10 $\mu\text{mol/mol}$) are generally more variable in shape, but still seem to be somewhat reproducible between shells. For the more precisely timed *P. maximus* data, the apexes of the large peaks all fall within 7 days of each other (this is a combined error of backdating striae, and time averaging several striae in each carbonate sample milled from the shell; Fig. 4). Nevertheless, the reproducibility between the shells strongly suggests that these peaks are a response to an exogenous factor in both species, or if not, the result of a very well-timed endogenous process.

There are several possible hypotheses that can be put forward regarding the cause of these Ba/Ca peaks. The link with primary production mentioned in the introduction above, variations in dissolved Ba/Ca in the water in which the bivalve is growing, variability in shell organic matter content, kinetic growth rate effects, and the consequence of a precisely timed and synchronous physiological process such as spawning are all plausible hypotheses. With the evidence at hand, we will discuss each of these possibilities.

In corals, Ba/Ca peaks have been shown to coincide with flood plumes that bring lower salinities and higher dissolved barium to the coastal zone (McCulloch et al.

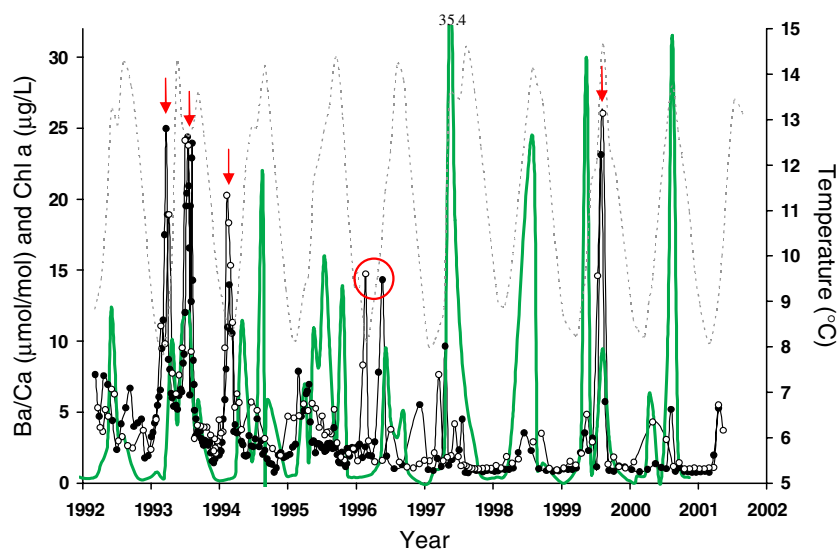


Fig. 2 Ba/Ca ratios in two *Saxidomus giganteus* shells that grew side by side in Puget Sound (solid and open circles). The solid green line is chlorophyll *a*, and the dotted line water temperature. Arrows indicate the synchronous sharp peaks noted in the text, and the red circle

highlights the mismatched section of the profile. Note that each year mark denotes the first day of that year. One chl *a* peak extends beyond the chart axis (labeled as 35.4 μg/l)

2003; Sinclair and McCulloch 2004). There is generally a linear inverse relationship between salinity and dissolved barium concentrations (Coffey et al. 1997; Sinclair and McCulloch 2004; Gillikin et al. 2006). Variability in dissolved barium concentrations as a source for the Ba/Ca peaks in *P. maximus* is highly unlikely, as the Bay of Brest is a rather stable environment with marine salinities year-round (Chauvaud et al. 2005). Additionally, if the background and peak incorporation mechanism were the same, the Ba/Ca ratio of the water would need to increase five to ten times, which is highly improbable considering salinity varies only a few units over the year (Lorrain et al. 2005). Salinity slowly increases from 33.1 to 35.6 from March to September, and can therefore not explain the Ba/Ca peaks measured in the shells (salinity mimics the temperature profile shown in Fig. 4; see Fig. 2 in Lorrain et al. 2005).

In bivalve shells, the Ba/Ca peaks are generally considered to be related to primary productivity (Stecher et al. 1996; Vander Putten et al. 2000; Lazareth et al. 2003). Phytoplankton decay, sedimentation, and release of barium should happen quickly after the end of a bloom in an estuary (Stecher and Kogut 1999; Ganeshram et al. 2003). Most of the barium released after a bloom is labile, and only a minor fraction forms barite crystals (Ganeshram et al. 2003). Therefore, if labile barium, either in phytoplankton or released into the dissolve phase, was the cause of the Ba/Ca peaks, these peaks should form near the end of the bloom, or very shortly thereafter. The Ba/Ca peaks are not systematically occurring near the end of the chl *a* peaks in *S. giganteus* shells; however, due to possible errors in aligning the shell and environmental data, we cannot conclude a lack of correlation. Nevertheless, it is clear that

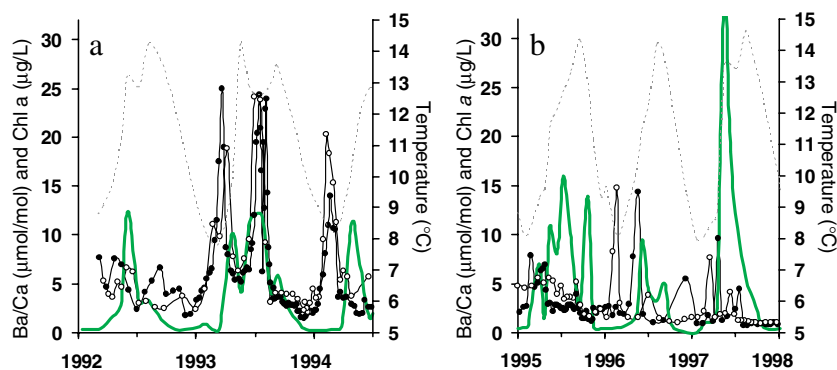


Fig. 3 An enlargement of two sections of the *Saxidomus giganteus* data presented in Fig. 2. The left panel (a) shows the first three large Ba/Ca peaks that correlate very well between the two shells, and the right panel (b) shows the section of shell with the poorest match

between the two shells. The difference in background levels though ontogeny is also more visible in these panels (i.e., they are higher in the right panel)

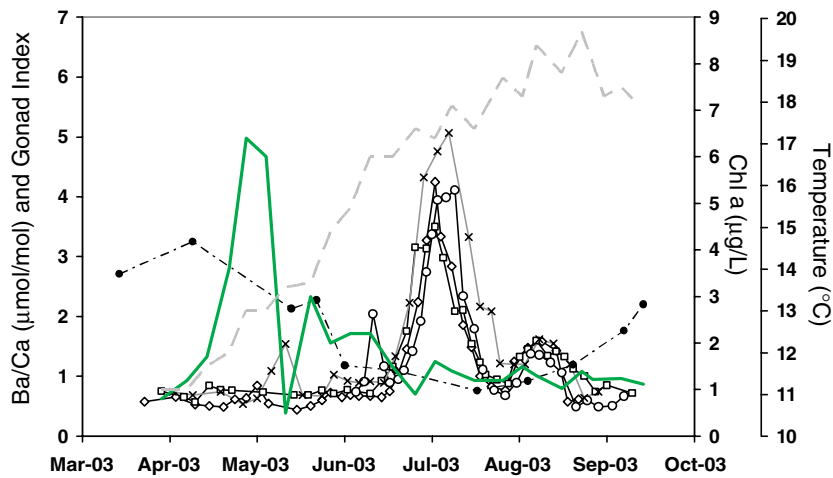


Fig. 4 Ba/Ca ratios in three age class I (open symbols) and one age class II (crosses) *Pecten maximus* shells that grew in the Bay of Brest (France). The solid green line is chlorophyll *a*, and the dashed grey line water temperature. The dashed black line and solid black symbols

are the gonad index (dry weight of the gonad divided by the shell weight, $\times 1,000$). Note that each month mark denotes the first day of that month

chl *a* amplitude is not related to Ba/Ca peak amplitude (also noted by Vander Putten et al. 2000 in *Mytilus edulis*); despite the large chl *a* peaks in 1997, 1998, and 2000, there are only minor Ba/Ca peaks during these years (Fig. 2). Considering the errors associated with the fit between shell and environmental data, we rely more heavily on the *P. maximus* shells, which are much more precisely timed.

The four *P. maximus* shells also show a high correlation between Ba/Ca peaks (Fig. 4). All shells exhibit a double peak, with the first large peak and second smaller peak synchronizing between all shells including both age classes, which illustrates that there is no ontogenic effect on these Ba/Ca profiles in *P. maximus*. There is a remarkable similarity in the chl *a* and Ba/Ca peak shapes (i.e., the presence of a large peak, followed by a smaller peak);

however, they are offset by about 40 days. The phytoplankton bloom crashed on the 12th of May, while the Ba/Ca peak started about the 21st of June. Moreover, the sedimentation of phytoplankton occurs around the end of the bloom, as is evidenced by a sharp reduction in growth rate (Fig. 5). This phenomenon has been previously described in Chauvaud et al. (1998) and in Lorrain et al. (2000); briefly, the growth reduction is caused by the sedimentation clogging the gills of the scallops and/or a reduction in oxygen at the water–sediment interface as a result of the decaying phytoplankton. Barite formation in settling phytoplankton flocs, or the Bar-rich phytoplankton itself, is clearly not the cause of these peaks, since the growth reduction and Ba/Ca peaks should then have coincided, which is not observed (Figs. 4 and 5). However, Ganeshram et al. (2003) found that barite

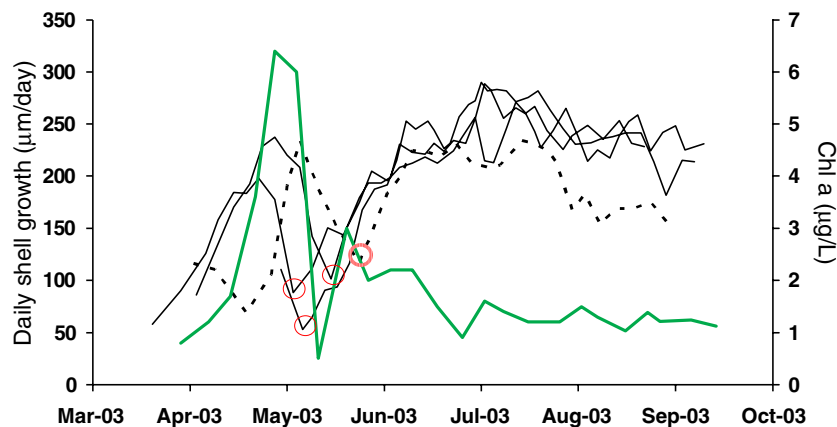


Fig. 5 Daily shell growth rates in three age class I (solid lines) and one age class II (dashed line) *Pecten maximus* shells that grew in the Bay of Brest (France). The solid green line is chlorophyll *a*. The red

circles highlight the growth minima during the growth dip. Note that each month mark denotes the first day of that month

formation can take several weeks to reach its maximum after the decay of phytoplankton is initiated. Perhaps barite is formed at the sediment surface, and is ingested by the scallops several weeks after the phytoplankton bloom ends. Although only a small amount of barium is converted to barite, the barite could provide a high concentration of Ba if it dissociates in the gut of the bivalve. It is convenient to suggest that there is indeed a link between phytoplankton and Ba/Ca peaks due to the synchronicity of the *S. giganteus* Ba/Ca profile, and the similarity in the peak shapes in *P. maximus*. However, it would be surprising that the double peak is persevered in *P. maximus*. If there was a biological filtering or delay in barite formation in the sediments, the signal in the shell would most likely be averaged, and would not show the discreet double peak. Detailed time-resolved experiments are needed to determine if this is the case. If barite is the cause, then it should be visible in the gut of bivalves at the time of Ba/Ca peak formation. Experiments over several years following the complete Ba cycle in the water and in the bivalve (e.g., different tissues, gut, and shell) are needed.

An alternative hypothesis could be that Ba residing in the animal's tissues is remobilized by the bivalve during spawning, as was proposed by Gillikin et al. (2006); however, *P. maximus* spawns at the beginning of the spring before the shell Ba/Ca peak, as attested by the gonadosomatic index (Fig. 4; Paulet et al. 1997). It is possible that Ba is used for reproduction, and is sequestered to the shell during non-reproductive periods, which would result in low shell Ba/Ca during reproduction and high shell Ba/Ca during non-reproductive periods, as seen here. Preliminary data show that gonads are rich in Ba when compared with muscle tissues (Lorrain, unpublished data). The fact that shell Ba/Ca peaks exactly when gonad development is at its lowest (Fig. 4) gives support to a Ba–reproduction link hypothesis. However, it seems unlikely that this would lead to similar variations in shell Ba/Ca peak amplitudes between individuals (Fig. 2). This again illustrates the need for further research involving the relationship between ecophysiology and proxy incorporation.

There are other potential causes of the Ba/Ca peaks in these shells. For example, if Ba were to bind to organic matter in the shell, then an enrichment in shell organic matter would result in a high Ba/Ca ratio. However, neither Hart et al. (1997) nor Sinclair (2005) found a relationship between organic matter and Ba concentrations in other biogenic carbonates (i.e., corals), and Rosenthal and Katz (1989) suggest that Ba is bound to the crystal phase, rather than the organic phase in mollusk shells. Moreover, the slower growing tip of the *S. giganteus* shells would be expected to have higher organic matter concentrations (Rosenberg et al. 2001), while higher Ba/Ca ratios are generally observed in the faster growing back sections of the shells (Fig. 2). Thus, it is unlikely that

the Ba peaks are associated with shell regions containing higher organic content. There has been one report of barite crystals in bivalve shells, but these shells came from a freshwater environment with extremely high barium concentrations (Fritz et al. 1990), so this is also an unlikely cause of the peaks.

Alternatively, the Ba/Ca peaks could be a kinetic growth rate effect. Such effects have been noted in both abiogenic calcite and aragonite (Tesoriero and Pankow 1996; Gaetani and Cohen 2006), but this is also not supported by our data. The double peak in *P. maximus* (Fig. 4) cannot be explained in terms of changing growth rates, as daily growth rates are fairly constant during peak formation (Fig. 5). Moreover, despite growth slowing from about 18 mm/year in 1993 to about 4 mm/year in 1999 in the *S. giganteus* shells (see Gillikin et al. 2005a), there are still large Ba/Ca peaks in these years (up to ~25 $\mu\text{mol/mol}$; Fig. 2). Therefore, there is no satisfactory variable that can be used to explain these ubiquitous Ba/Ca peaks in bivalve shells. Interestingly, Sinclair (2005) also found well-synchronized Ba/Ca peaks in corals that could not be correlated to several environmental and biological processes.

Ba partition coefficients and shell Ba/Ca background

The partitioning of trace elements (Me) between water and carbonate is typically defined in terms of a partition coefficient: $D_{\text{Me}} = (\text{Me/Ca})_{\text{carbonate}}/(\text{Me/Ca})_{\text{water}}$, where Me/Ca are typically given as molar ratios (see Mucci and Morse 1990). Interestingly, abiogenic calcite and aragonite behave rather differently with regards to Ba partitioning (D_{Ba}). Typically, abiogenic calcite D_{Ba} is precipitation rate-dependent, and is low, usually <0.5 (Mucci and Morse 1990; Tesoriero and Pankow 1996). Abiogenic aragonite D_{Ba} is strongly temperature-dependent, and is usually >0.5 (Dietzel et al. 2004; Gaetani and Cohen 2006). However, both the aragonite (*S. giganteus*) and calcite (*P. maximus*) shells presented here have similar baseline values (*S. giganteus* ~1.0, and *P. maximus* ~0.7 $\mu\text{mol/mol}$). Considering a seawater Ba/Ca ratio of 3.8 $\mu\text{mol/mol}$ for the Bay of Brest (Lorrain, unpublished data), and a background Ba/Ca ratio of 0.7 $\mu\text{mol/mol}$ for *P. maximus*, a D_{Ba} of 0.18 can be calculated. Similarly, *S. giganteus* has a D_{Ba} of 0.16 using Ba/Ca ratios of 1.0 $\mu\text{mol/mol}$ for the shell, and 6.2 $\mu\text{mol/mol}$ for the water (Gillikin 2005). These are both close to the D_{Ba} that Gillikin et al. (2006) report for *M. edulis* (~0.1). In addition, similar partition coefficients for both mineralogies are evident from the peaks. Aragonite shells had higher peaks in this study, but previous studies presented even higher peaks in calcite shells (e.g., *M. edulis* up to 70 $\mu\text{mol/mol}$; Vander Putten et al. 2000). Thus, despite different mineralogies, these species have very similar barium partition coefficients.

While most research has focused on the Ba peaks, Gillikin et al. (2006) found that the background Ba/Ca level in the shells of another bivalve species (*Mytilus edulis*, calcite shell) could be used as an indicator of dissolved Ba/Ca ratios in estuarine waters, and hence provide an indication of salinity (cf. salinity varies inversely with dissolved Ba/Ca ratios in most estuaries). However, in the two *S. giganteus* shells analyzed here, the background Ba/Ca levels are higher near the umbo (Fig. 2). Considering that the $\delta^{18}\text{O}$ data do not suggest a reduced salinity when the clams were younger (see Gillikin et al. 2005b), and the salinity record does not indicate reduced salinities in the Sound, this is probably an ontogenic effect. Unfortunately, this complicates the use of Ba/Ca backgrounds in this species, but does highlight the point that species-specific differences in elemental incorporation are important, and must be taken into consideration.

Despite the ontogenic effect, large temporary salinity changes may still be recorded in this species' shells, especially if shell regions of similar age are compared. Considering that the D_{Ba} for aragonitic corals (1.27; Lea et al. 1989) and aragonitic coralline sponges (0.86; Rosenheim et al. 2005) is much closer to the abiogenic end-member, it can be postulated that these aragonitic bivalves are strongly regulating Ba incorporation into their shells. This is also evidenced by the fact that Dietzel et al. (2004) report aragonite D_{Ba} to be 2.0 at 10°C, and 1.5 at 25°C. This temperature dependence in aragonite has also been demonstrated in larval mollusks (Zacherl et al. 2003), coralline sponges (Rosenheim et al. 2005), and corals (Lea et al. 1989), while it is obviously not evident in the shells used in this study (if it were, we would expect to see a sinusoidal curve in the *S. giganteus* data). Therefore, barium incorporation into bivalve carbonate is not similar to abiogenic carbonate, but does still seem to record some environmental signal.

In conclusion, this study suggests that Ba/Ca peaks in bivalve shells are caused by an environmental forcing regardless of mineralogy, albeit with a biological filter. The exact timing between animals growing in the same environment attests to this. Clearly, more work is needed to decipher this potential proxy.

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References

- Bishop JKB (1988) The barite–opal–organic carbon association in oceanic particulate matter. *Nature* 332:341–343
- Böhm F, Joachimski MM, Dullo WC, Eisenhauer A, Lehnert H, Reitner J, Worheide G (2000) Oxygen isotope fractionation in marine aragonite of coralline sponges. *Geochim Cosmochim Acta* 64:1695–1703
- Carré M, Bentaleb I, Bruguier O, Ordinola E, Fontugne M (2006) Calcification rate influence on trace elements incorporation in marine bivalve aragonite: evidences and mechanisms. *Geochim Cosmochim Acta* 70:4906–4920
- Chauvaud L, Thouzeau G, Paulet Y-M (1998) Effects of environmental factors on the daily growth rate of *Pecten maximus* juveniles in the Bay of Brest (France). *J Exp Mar Biol Ecol* 227:83–111
- Chauvaud L, Lorrain A, Dunbar RB, Paulet Y-M, Thouzeau G, Jean F, Guarini J-M, Mucciarone D (2005) The shell of the great scallop *Pecten maximus* as a high frequency archive of paleoenvironmental change. *Geochem Geophys Geosyst* 6:Q08001. doi:10.1029/2004GC000890
- Coffey M, Dehairs F, Collette O, Luther G, Church T, Jickells T (1997) The behaviour of dissolved barium in estuaries. *Estuar Coast Shelf Sci* 45:113–121
- Dehairs F, Chesselet R, Jedwab J (1980) Discrete suspended particles of barite and the barium cycle in the open ocean. *Earth Planet Sci Lett* 49:528–550
- De Ridder F, Pintelon R, Schoukens J, Gillikin DP, André L, Baeyens W, de Brauwere A, Dehairs F (2004) Decoding nonlinear growth rates in biogenic environmental archives. *Geochem Geophys Geosyst* 5:Q12015. doi:10.1029/2004GC000771
- Dietzel MN, Gussone N, Eisenhauer A (2004) Co-precipitation of Sr^{2+} and Ba^{2+} with aragonite by membrane diffusion of CO_2 between 10 and 50 degrees C. *Chem Geol* 203:139–151
- Freitas PS, Clarke LJ, Kennedy H, Richardson CA, Abrantes F (2006) Environmental and biological controls on elemental (Mg/Ca, Sr/Ca and Mn/Ca) ratios in shells of the king scallop *Pecten maximus*. *Geochim Cosmochim Acta* 70:5119–5133
- Fritz LW, Ragone LM, Lutz RA (1990) Biomineralization of barite in the shell of the freshwater Asiatic clam *Corbicula fluminea* (Mollusca: Bivalvia). *Limnol Oceanogr* 35:756–762
- Gaetani GA, Cohen AL (2006) Element partitioning during precipitation of aragonite from seawater: a framework for understanding paleoproxies. *Geochim Cosmochim Acta* 70:4617–4634
- Ganeshram RS, Francois R, Commeau J, Brown-Leger SL (2003) An experimental investigation of barite formation in seawater. *Geochim Cosmochim Acta* 67:2599–2605
- Gillikin DP (2005) Geochemistry of marine bivalve shells: the potential for paleoenvironmental reconstruction. PhD Thesis, Vrije Universiteit Brussel, Belgium
- Gillikin DP, Lorrain A, Navez J, Taylor JW, André L, Keppens E, Baeyens W, Dehairs F (2005a) Strong biological controls on Sr/Ca ratios in aragonitic marine bivalve shells. *Geochem Geophys Geosyst* 6:Q05009. doi:10.1029/2004GC000874
- Gillikin DP, De Ridder F, Ulens H, Elskens M, Keppens E, Baeyens W, Dehairs F (2005b) Assessing the reproducibility and reliability of estuarine bivalve shells (*Saxidomus giganteus*) for sea surface temperature reconstruction: implications for paleoclimate studies. *Palaeogeogr Palaeoclimatol Palaeoecol* 228:70–85
- Gillikin DP, Dehairs F, Lorrain A, Steenmans D, Baeyens W, André L (2006) Barium uptake into the shells of the common mussel (*Mytilus edulis*) and the potential for estuarine paleo-chemistry reconstruction. *Geochim Cosmochim Acta* 70:395–407
- Hart SR, Cohen AL, Ramsay P (1997) Microscale analysis of Sr/Ca and Ba/Ca in Porites. In: *Proc 8th Int Coral Reef Symp* 2:1707–1712

- Kingston AW, Gröcke DR, Burchell M (2008) A multi-axial growth analysis of stable isotopes in the modern shell of *Saxidomus gigantea*: implications for sclerochronology studies. *Geochem Geophys Geosyst* 9:Q01007. doi:10.1029/2007GC001807
- Lazareth CE, Vander Putten E, André L, Dehairs F (2003) High-resolution trace element profiles in shells of the mangrove bivalve *Isognomon ephippium*: a record of environmental spatio-temporal variations? *Estuar Coast Shelf Sci* 57:1103–1114
- Lea DW (2003) Elemental and isotopic proxies of marine temperatures. In: Elderfield H (ed) *The oceans and marine geochemistry*. Elsevier-Perгамon, Oxford, pp 365–390
- Lea DW, Shen GT, Boyle EA (1989) Coralline barium records temporal variability in equatorial Pacific upwelling. *Nature* 340:373–376
- Lorrain A, Paulet Y-M, Chauvaud L, Savoye N, Nézan E, Guérin L (2000) Growth anomalies in *Pecten maximus* from coastal waters (Bay of Brest, France): relationship with diatom blooms. *J Mar Biol Assoc UK* 80:667–673
- Lorrain A, Paulet Y-M, Chauvaud L, Dunbar R, Mucciarone D, Fontugne M (2004) $\delta^{13}\text{C}$ variation in scallop shells: increasing metabolic carbon contribution with body size? *Geochim Cosmochim Acta* 68:3509–3519
- Lorrain A, Gillikin DP, Paulet Y-M, Chauvaud L, Le Mercier A, Navez J, André L (2005) Strong kinetic effects on Sr/Ca ratios in the calcitic bivalve *Pecten maximus*. *Geology* 33:965–968
- McCulloch MT, Fallon S, Wyndham T, Hendy E, Lough J, Barnes D (2003) Coral record of increased sediment flux to the inner Great Barrier Reef since European settlement. *Nature* 421:727–730
- Mucci A, Morse JW (1990) Chemistry of low-temperature abiotic calcites: experimental studies on coprecipitation, stability and fractionation. *Rev Aquat Sci* 3:217–254
- Paulet Y-M, Fifas S (1989) Potential fecundity of the scallop, *Pecten maximus*, from the bay of Saint-Brieuc. *Haliotis* 19:275–285
- Paulet Y-M, Dorange G, Cochard JC, Le Pennec M (1992) Reproduction and recruitment in *Pecten maximus* L. *Ann Inst Océanogr Paris* 68:1:45–64
- Paulet Y-M, Bekhadra F, Dechauvellerie N, Donval A, Dorange G (1997) Seasonal cycles, reproduction and oocyte quality in *Pecten maximus* from the Bay of Brest. *Ann Inst Océanogr Paris* 73:101–112
- Pearce NJG, Perkins WT, Westgate JA, Gorton MP, Jackson SE, Neal CR, Chenery SP (1997) A compilation of new and published major and trace element data for NIST SRM 610 and NIST SRM 612 glass reference materials. *Geostand Newslett* 21:115–144
- Purton LMA, Shields GA, Brasier MD, Grime GW (1999) Metabolism controls Sr/Ca ratios in fossil aragonitic mollusks. *Geology* 27:1083–1086
- Rosenberg GD, Hughes WW, Parker DL, Ray BD (2001) The geometry of bivalve shell chemistry and mantle metabolism. *Am Malacol Bull* 16:251–261
- Rosenheim BE, Swart PK, Thorrold SR (2005) Minor and trace elements in sclerosponge *Ceratoporella nicholsoni*: biogenic aragonite near the inorganic endmember? *Palaeogeogr Palaeoclimatol Palaeoecol* 228:109–129
- Rosenthal Y, Katz A (1989) The applicability of trace elements in freshwater shells for paleochemical studies. *Chem Geol* 78:65–76
- Sinclair DJ (2005) Non river-flood barium signals in the skeletons of corals from coastal Queensland, Australia. *Earth Planet Sci Lett* 237:354–369
- Sinclair DJ, McCulloch MT (2004) Corals record low mobile barium concentrations in the Burdekin River during the 1974 flood: evidence for limited Ba supply to rivers? *Palaeogeogr Palaeoclimatol Palaeoecol* 214:155–174
- Stecher HA, Kogut MB (1999) Rapid barium removal in the Delaware estuary. *Geochim Cosmochim Acta* 63:1003–1012
- Stecher HA, Krantz DE, Lord CJ, Luther GW, Bock KW (1996) Profiles of strontium and barium in *Mercenaria mercenaria* and *Spisula solidissima* shells. *Geochim Cosmochim Acta* 60:3445–3456
- Surge D, Walker KJ (2006) Geochemical variation in microstructural shell layers of the southern quahog (*Mercenaria campechiensis*): implications for reconstructing seasonality. *Palaeogeogr Palaeoclimatol Palaeoecol* 237:182–190
- Takesue RK, van Geen A (2004) Mg/Ca, Sr/Ca, and stable isotopes in modern and Holocene *Protothaca staminea* shells from a northern California coastal upwelling region. *Geochim Cosmochim Acta* 68:3845–3861
- Tesoriero AJ, Pankow JF (1996) Solid solution partitioning of Sr^{2+} , Ba^{2+} , and Cd^{2+} to calcite. *Geochim Cosmochim Acta* 60:1053–1063
- Vander Putten E, Dehairs F, Keppens E, Baeyens W (2000) High resolution distribution of trace elements in the calcite shell layer of modern *Mytilus edulis*: environmental and biological controls. *Geochim Cosmochim Acta* 64:997–1011
- Walter E, Pronzato L (1994) Identification de modèles paramétriques à partir de données expérimentales. Masson, Paris
- Zacherl DC, Paradis G, Lea DW (2003) Barium and strontium uptake into larval protoconchs and statoliths of the marine neogastropod *Kelletia kelledi*. *Geochim Cosmochim Acta* 67:4091–4099